

Atmospheric phenomena deduced from radiosonde and GPS occultation measurements for various application related studies

C J JOHNY^{1,*}, S K SARKAR^{1,**} and D PUNYASESUDU^{2,†}

¹*Radio & Atmospheric Sciences Division, National Physical Laboratory,
Dr. K S Krishnan Marg, New Delhi 110 012, India.*

²*Department of Physics, Post Graduate Study Centre, Sri Krishnadevaraya University,
Kurnool 518 002, A.P., India.*

**e-mail: johnycj@rediffmail.com*

***e-mail: sksarkar@mail.nplindia.ernet.in*

†e-mail: dupadu@rediffmail.com

The tropopause height and tropopause temperature are sensitive to temperature changes in troposphere and stratosphere. These are the measures of global climatic variability. Atmospheric profiles of temperature, refractivity and water vapour are always needed for communication, navigation and atmospheric modeling studies. The tropopause characteristics over the Indian region have been studied using radio occultation measurements (CHAMP) on the basis of cold point criterion. Tropopause height shows large variation in the latitude range $\sim 30^{\circ}$ – 40° N during winter. Tropopause temperature less than -82°C , assumed to facilitate troposphere to stratosphere air transport, is observed at a number of tropical Indian locations and no seasonal pattern is observed in its occurrence. The bias in temperature and refractivity deduced from radiosonde and radio occultation measurements is also presented.

1. Introduction

Troposphere is the lowest portion of earth's atmosphere and hold approximately 80% of atmospheric airmass. The radio wave propagation through troposphere depends on refractive index of the medium. The refractive index of the medium depends on atmospheric pressure, temperature and humidity. The properties of radio wave make it an efficient tool to study the atmosphere. Earlier, atmospheric monitoring was based on radiosonde measurements. The availability of global positioning system (GPS) signals has given a new technique for studying the atmosphere. The GPS satellites orbiting around 20,000 km above the surface of the earth continuously emits radio signals. The GPS signals received on low earth orbiting satellite

(LEO) is getting occulted by atmosphere as GPS satellite rises above or sets behind the horizon of the earth. Atmosphere causes refraction of radio signals as it passes through it. This time delay and bending is the basis of estimation of atmospheric refractivity and temperature (Wickert *et al* 2001). Radio occultation technique is free of interference from cloud, aerosol, precipitation and it provides global coverage. Radio occultation measurements combined with radiosonde measurements give a better picture of atmospheric activities.

There were some comparison studies of radio occultation observations with radiosonde observations and meteorological analyses from weather models. Kurisinky *et al* (1997) reported mean bias of 0.5°C between radio occultation and European Centre for Medium Range Weather

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Forecasts (ECMWF) weather analyses over northern hemisphere. Wickert *et al* (2004) reported radio occultation observations show a mean temperature bias of less than 0.4°C between meteorological analyses from ECMWF in the altitude range 10–25 km based on comparison studies performed during the period from March 2002 to December 2002. They confirmed the results with global radiosonde network. About 69% of the radiosonde observations used for comparison studies are in the range $30\text{--}70^{\circ}\text{N}$. In tropical region, temperature bias in relation to analyses is found to be larger than that with radiosonde observations. A temperature bias of less than 1.5°C in the pressure range 300–10 mb (Schmidt *et al* 2005) is reported between radio occultation and radiosonde observations. These results are based on the observations between May 2001 and December 2004. The comparison of two CHAMP temperature profiles over Sepang (2.71°N , 101.7°E) and Singapore (1.36°N , 103.98°E) with radiosonde and analyses from National Centers for Environmental Prediction (NCEP) showed excellent agreement over all heights below tropopause (Kishore *et al* 2006). Only very few comparative studies of radio occultation and radiosonde observations were reported over the Indian region. Kuo *et al* (2005) made comparison of vertical refractivity profiles from radio occultation measurements with the radiosonde observations during the period from June 2001 to March 2004 over India, China, Japan, Australia and Russia.

The radio refractivity data is required for design of tropospheric communication circuits and radar operations as well as remote sensing applications, which are widely used in defense and aviation sector. Temperature observations are needed in weather forecasts and global climate change related studies. The global coverage of occultation measurements makes it useful in the areas where radiosonde observations are limited, particularly in oceanic regions.

Tropopause is the boundary between turbulent troposphere and more stable stratosphere and plays a significant role in stratospheric tropospheric exchange. Tropopause affects the transport of ozone, water vapour and aerosol between troposphere and stratosphere and has an important role in many dynamic weather changes in troposphere like cyclones and jet streams. Coupling between troposphere and stratosphere by dynamical, chemical and radiative processes has an effect on climate variability and global change. It is found that wave-induced forces withdraw air upward from troposphere in tropics and pushes through lower stratosphere towards poles and pushes it downwards in extra tropics (Holton *et al* 1995). Different definitions exist for tropopause, based

on criteria like variation in temperature lapse rate, potential vorticity, minimum temperature and ozone gradient. Convective activities occurring in the troposphere have strong influence on tropical tropopause region. Convective processes are found to correlate better with minimum temperature or cold point tropopause (CPT) than other definitions of tropopause (Schmidt *et al* 2005). In studies related to tropospheric stratospheric exchange CPT definition is preferred. Tropical tropopause is colder and extremely drier than at higher latitudes (Jain *et al* 2006).

The tropopause structure has strong influence on tropospheric stratospheric exchange. It is considered that air enters from troposphere to stratosphere primarily through tropics. According to fountain hypothesis, tropospheric air enters stratosphere through areas where tropopause temperature is less than their mean longitudinal values to freeze dry air entering stratosphere (Zhou *et al* 2001). Stratospheric water vapour concentration is influenced by tropopause structure although exact relation between them is not clear (Schmidt *et al* 2005). Considering water vapour concentration in stratosphere, it is estimated that troposphere to stratosphere transport occurs through areas where temperature is less than -82°C (Zhou *et al* 2001). By studying temperature distribution, it is assumed that during winter season tropospheric air enters through Indonesian region (Newel and Gould 1981). It is found in some studies tropopause temperature as low as -82°C occurs in tropical Indian region during summer monsoon and post-monsoon months (Jain *et al* 2006). The occurrence of low tropopause temperature is found to be mostly in oceanic region compared to land (Jain *et al* 2006). The knowledge of tropopause structure and its variability are important in understanding processes involved in climate variability. On large time scales tropopause variation gives information about general circulation processes (Angell and Korshover 1974). In shorter scales phenomena like tropopause folds believed to play an important role in exchange of chemical species and vorticity.

The tropopause height is sensitive to temperature changes in the troposphere and stratosphere caused by both natural and human factors. Thus, the height of the tropopause provides another fingerprint of human effects on climate. It is lower at higher latitudes and higher at lower latitudes. In tropics it is above 16 km and in extra tropics it is between 8 and 12 km. The tropopause height variation is also one of the mechanisms through which the tropospheric–stratospheric exchange process can be understood. The stratospheric water vapour concentration is found to be increased in recent years. The source of water vapour in stratosphere

is from methane oxidation and transport from tropical troposphere. One possible mechanism for such increase through transport from tropical troposphere could be the temporal increase in tropical tropopause temperature or tropopause height, which could allow large saturation mixing ratios to enter stratosphere (Randel *et al* 2000).

The subtropics are the region of transition between tropics and extra tropics characterized by split in tropopause. Atmospheric phenomena like tropopause fold and subtropical jet stream occurs in this region. Jet streams are fast flowing narrow currents occurring in the atmosphere near tropopause. Stratospheric intrusion of air into the baroclinic zone beneath upper tropospheric jet stream is known as tropopause fold. Tropopause folds are a mechanism of stratospheric tropospheric exchange in which dry stratospheric air rich in ozone transported downwards to tropospheric levels and tropopause height drops suddenly as the stratospheric air replaced in fold is warmer than tropospheric air. Double tropopauses (lapse rate criteria) are found to occur frequently in this region during winter months associated with cyclonic circulation systems (Randel *et al* 2007). This may be due to above jet streams stability of lower stratosphere is reduced giving rise to poleward transport of upper tropospheric air. The actual mechanism behind the transport is not clear and it does not have any relation with the occurrence of storm tracks or Rossby wave breaking at tropopause, which occurs in summer season (Randel *et al* 2007). Extra tropics are characterized by large dynamic variability characterized by complex spatial structure.

The results on radio refractivity and temperature profile over two Indian tropical stations, Kolkata (22.65°N, 88.15°E) and New Delhi (28.58°N, 77.20°E) deduced from radiosonde and radio occultation measurements are presented in this paper. The tropopause characteristics over Indian region have also been studied using radio occultation (RO) observations.

2. Database

The radiosonde observations were obtained from British Atmospheric Data Center (BADC) and GPS/CHAMP (Challenging Mini Pay Load) radio occultation observations (GeoForschungs-Zentrum (GFZ), Potsdam, Germany) obtained through Information System Data Centre (ISDC) (<http://isdc.gfz-potsdam.de/champ>). The CHAMP/GPS satellite was launched on July 15, 2000 into an almost circular and near polar orbit (with the inclination of 87°) with an initial altitude of 454 km. Black Jack GPS receiver

developed by Jet Propulsion Laboratory capable of tracking GPS L2 (frequency, L2 ~ 1.277 GHz) signals modulated with Y code is used (Kuo *et al* 2004). The CHAMP mission provides observations at a vertical resolution of ~200 m at an average of ~260 atmospheric profiles per day. The occultation data processing system uses occultation data by CHAMP, ground tracking data of GPS satellites and precise orbit information of satellites. The phase observations made by CHAMP are synchronized with observations of reference ground network stations. Atmospheric bending angle is calculated from measured time series of phase path increase of two GPS frequencies (L1 ~ 1.575 GHz, L2) and of satellite velocities and positions. Atmospheric refractivity is derived from bending angles using Abel inversion. In dry conditions atmospheric pressure and temperature are deduced from refractivity using hydrostatic equation. In wet conditions, additional background information is utilized for deriving temperature. Radiosonde measurements are taken twice a day at 00:00 and 12:00 (UTC) pertaining to early morning and late evening Indian local time. IM-MK3 radiosonde is used for measurement (Kuo *et al* 2005). White painted rod thermistor is used for sensing temperature.

3. Comparison of CHAMP and radiosonde measurements

The comparison of CHAMP observations with radiosonde measurements is performed to examine the reliability of using both together for various applications. The radio refractivity and temperature observations at different altitude levels are provided directly by CHAMP. But, in case of radiosonde, radio refractivity is estimated from pressure, temperature and dew point temperature by using following relations.

Refractivity N is calculated using expression (1):

$$N = (77.6 P/T) + (3.73 \times 10^5 e/T^2), \quad (1)$$

where P is pressure in mb, T is temperature in Kelvin, e is vapour pressure in mb.

Vapour pressure is calculated from dew point temperature using expression (2):

$$e = es_0 \times \exp \left(\frac{lv}{Rv} \times \left(\frac{1}{T_0} - \frac{1}{T_d} \right) \right), \quad (2)$$

where es_0 is the reference saturation vapour pressure at STP = 6.11 mb; T_0 is the reference

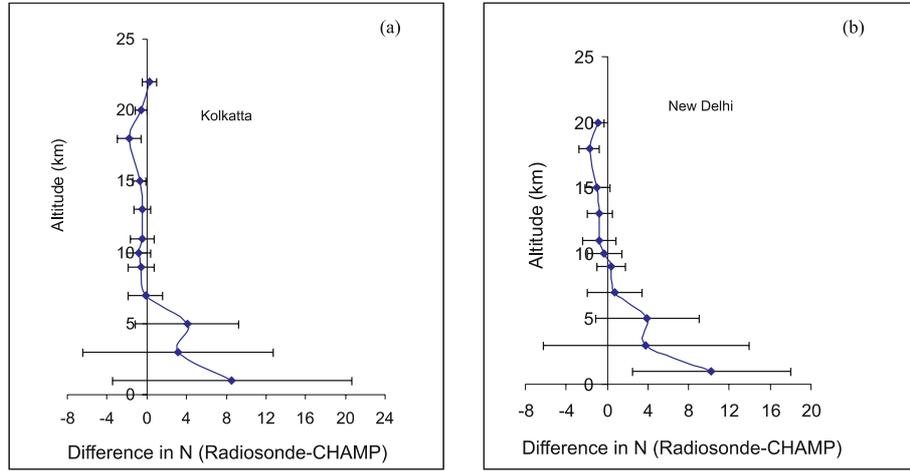


Figure 1. Profile of the mean value of the difference in refractivities between radiosonde and CHAMP data at (a) Kolkata and (b) New Delhi stations. (Horizontal bars indicate the standard deviation of the data.)

temperature at STP = 273.15 K; T_d is dew point temperature (K); T is temperature (K); lv is latent heat of vapourization of water = 2.5×10^6 joules per kilogram; and Rv is gas constant for water vapour (461.5 joules Kelvin/kilogram).

The results on radio refractivity and temperature deduced from the radiosonde measurements over Kolkata and New Delhi have been compared with CHAMP observations at different altitude levels. The comparison is performed with the criteria that both measurements are within the range of two degrees of position and time difference between measurements is less than 5 hours. The analysis is done for the period from 2002 January 1 to 2006 December 31. The CHAMP profiles in the range (lat. 20° – 24° N, long. 86° – 90° E) are considered for Kolkata and in the range (lat. 27° – 31° N, long. 75° – 79° E) for New Delhi.

4. Tropopause structure measurement

The tropopause height and temperature is deduced for the latitude range 2° – 40° and longitude range 65° – 100° for the time period of 2002 to 2006 from radio occultation observations. The tropopause is detected using cold point criterion. The spatial and temporal coverage of occultation measurements are different for different years. This is overcome by estimating annual average of tropopause height and temperature at one degree spatial interval for the latitude and longitude range mentioned above.

5. Comparison of vertical atmospheric profiles

A total of 23 profiles obtained for Kolkata and 31 profiles over New Delhi have been utilized for

comparison purpose. In lower atmospheric layers below 4 km radio occultation measurements provide only few observations. The radio refractivity measurements show large variations below 5 km. The mean value of difference in refractivity between CHAMP and radiosonde measurements (CHAMP-radiosonde) is deduced for different altitudes and plotted against altitude. The results of radio refractivity are shown in figure 1. The horizontal bars indicate standard deviation of radio refractivity. The difference in refractivity between both modes of measurements is large below 5 km altitude and above 5 km, the difference in refractivity between two modes of measurements is less than $2N$. The variation between the results on radio refractivity from radiosonde and radio occultation in lower troposphere may be due to the large scale variation of water vapour content in the Indian region. The humidity in lower troposphere over tropics is significant.

The temperature measurements also show large variation between radio occultation and radiosonde observations below 5 km altitude. Figure 2 shows the comparison of typical temperature and refractivity profiles observed on a particular day. Above 5 km altitude, bias between measurements decreases and above 10 km altitude bias reduces considerably. In the altitude range from 10 km to tropopause altitude, the mean value of bias between the measurements is less than 2.5°C . The radio refractivity derived from GPS occultation measurement is composed of dry and wet contribution. In deriving temperature and pressure from refractivity profiles, information about water vapour is required (Kurisinky *et al* 1997). The required background information of humidity is taken from operational meteorological analysis of European Centre for Medium-Range Weather

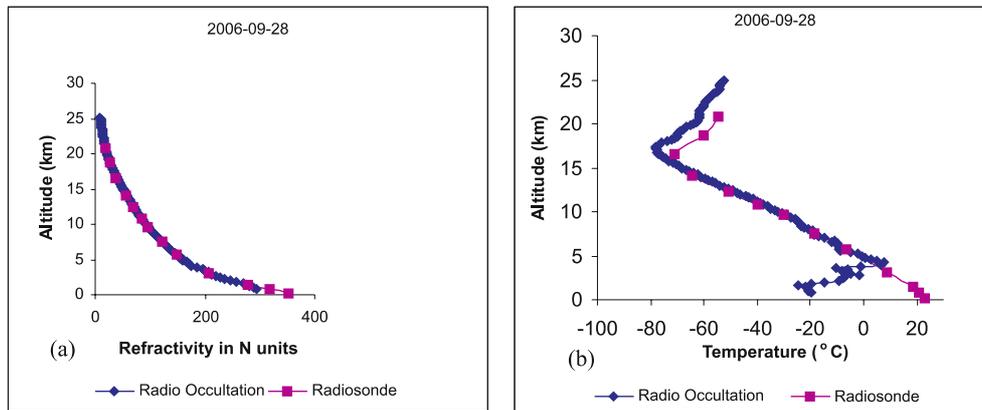


Figure 2. Typical vertical profile of (a) refractivity and (b) temperature obtained with CHAMP and radiosonde observations at New Delhi observed on a particular day.

Forecasts (ECMWF) (Wickert *et al* 2004). Dry atmosphere and hydrostatic equilibrium is assumed above 10–12 km in deducing atmospheric parameters from bending angle (Narayana Rao *et al* 2007). These assumptions may not be strictly valid during strong convections and presence of water vapour in upper troposphere. In tropical tropopause region strong convections and cirrus clouds are reported to occur frequently (Sunil Kumar *et al* 2003; Narayana Rao *et al* 2007). It is reported in tropical upper troposphere, ECMWF analyses differ from real atmospheric situations (Wickert *et al* 2004). The comparison studies temperature measurements of different type of radiosondes with radio occultation observations reported different responses (Kuo *et al* 2005). The accuracy of radiosonde measurements depends on the temperature sensor error and error in height (pressure) measurements. Solar heating produces significant errors in temperature measurements below 100 mb pressure levels and in cloudy conditions.

6. Tropopause height

The results on tropopause height are based on 2729 occultation profiles during the period. The tropopause height is found to vary between 9 and 23 km and temperature between -57° and -96°C . The mean value of tropopause height is found to be 17.3 km that is well in the range of earlier observations (Venkat Ratnam *et al* 2005). In 90% of the observations tropopause height is in the range of 16–19 km. The histogram of the tropopause height with latitude is given in figure 3. The high tropopause height characteristic of tropical region is found to exist up to 40°N . There are many studies suggesting the widening of tropical region towards poles as a result of jet

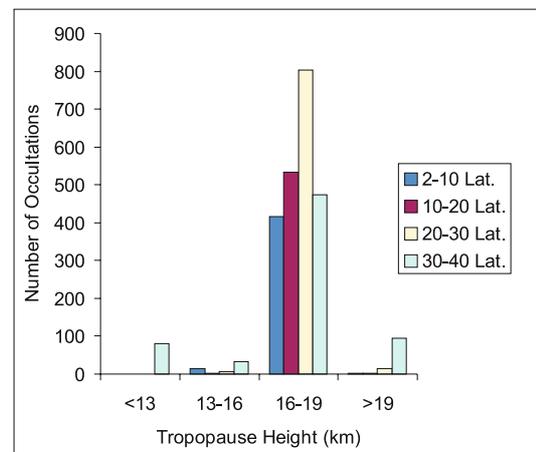


Figure 3. Histogram analysis of tropopause height for different latitude ranges.

streams, circulation systems and anthropogenic activities (Fu *et al* 2006; Hudson *et al* 2006). The expansion of tropical tropopause around 5° to 8° towards poles is reported in the radiosonde studies conducted in the period of 1979–2005 (Seidel and Randel 2007). The tropopause height in the tropical region is found to be within 15–20 km. The contour maps showing geographical variation of tropopause height and tropopause temperature are shown in figures 4 and 5 respectively.

It is seen that tropopause height shows small variations meridionally and zonally in the latitude range 2° – 30°N . But, in the range of 30° – 40°N latitude large variation in tropopause height occur zonally. The tropopause temperature is found to increase with latitude and higher tropopause temperature is observed in the latitude range of 30° – 40°N . Figure 6 shows the meridional variation of tropopause temperature in this period. It is also reported earlier that tropopause temperature increases with latitude (Kishore *et al* 2006).

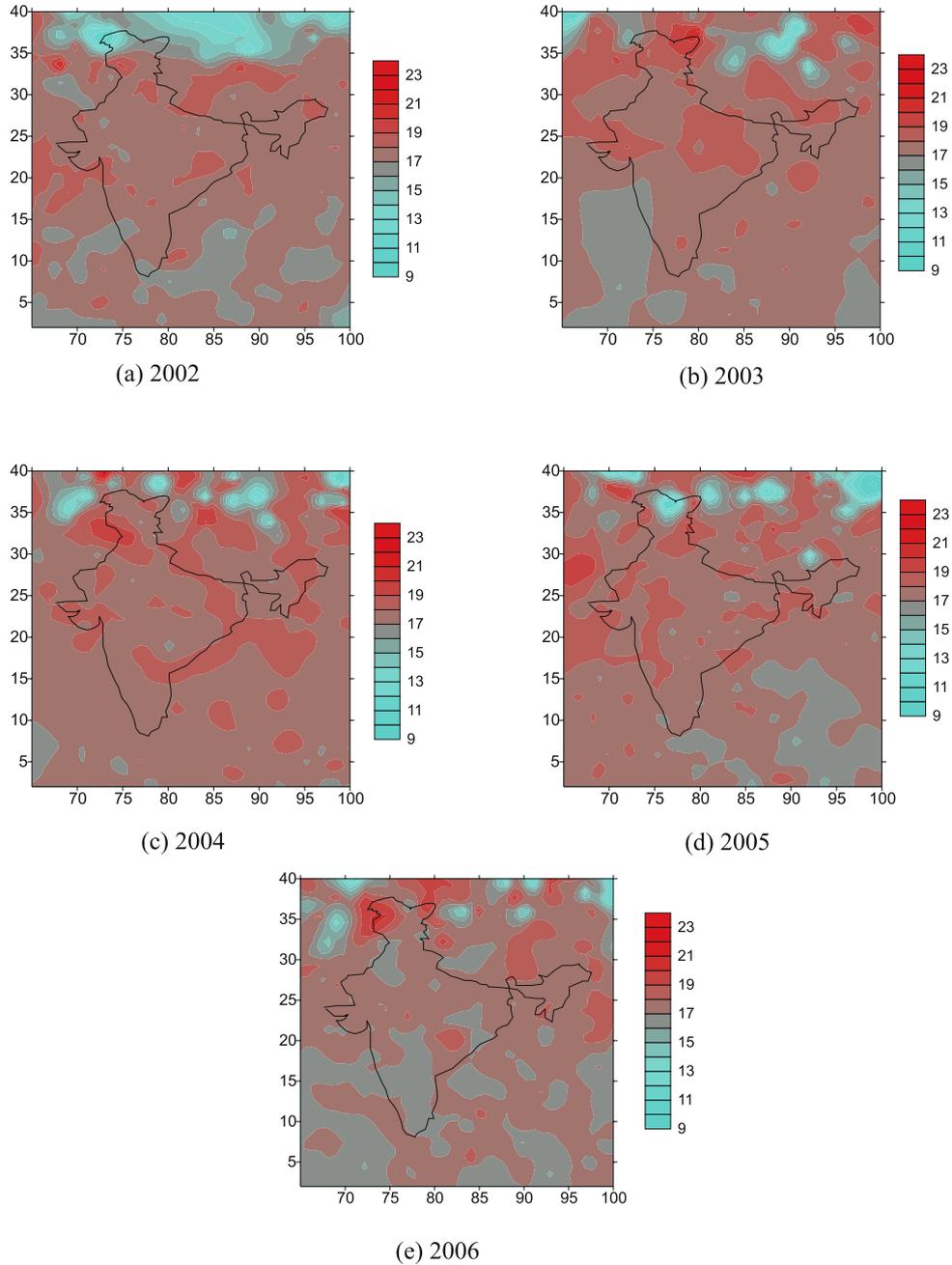


Figure 4. Geographical variation of tropopause height in (km) for different years.

7. Cold tropopause temperature

Out of all observations pertaining to the period from 2002 to 2006 tropopause temperature less than -82°C (cold tropopause temperature) is observed on 24% of time. Figure 7 shows year-wise distribution of cold tropopause temperature along with observation points. The cold tropopause temperature is not obtained beyond 32°N . In the year 2002, the number of cold tropopauses obtained above 23.5°N latitude is more than that in other years. Considering tropical region alone, cold tropopause was obtained for 44% of the cases, while

in 2006, in tropical region cold tropopause was obtained only in 30% of the cases. The tropopause height corresponding to cold tropopause varies between 16 and 19.6 km. The present observations confirm the previous studies made in the region using radiosonde measurements during monsoon and post-monsoon months about the existence of cold tropopauses (Jain *et al* 2006). No seasonal or diurnal bias is observed in the occurrence of these low tropopause temperatures. There were studies, which also reported mean monthly tropopause temperature less than -82°C in the tropical Pacific region during the months November to April

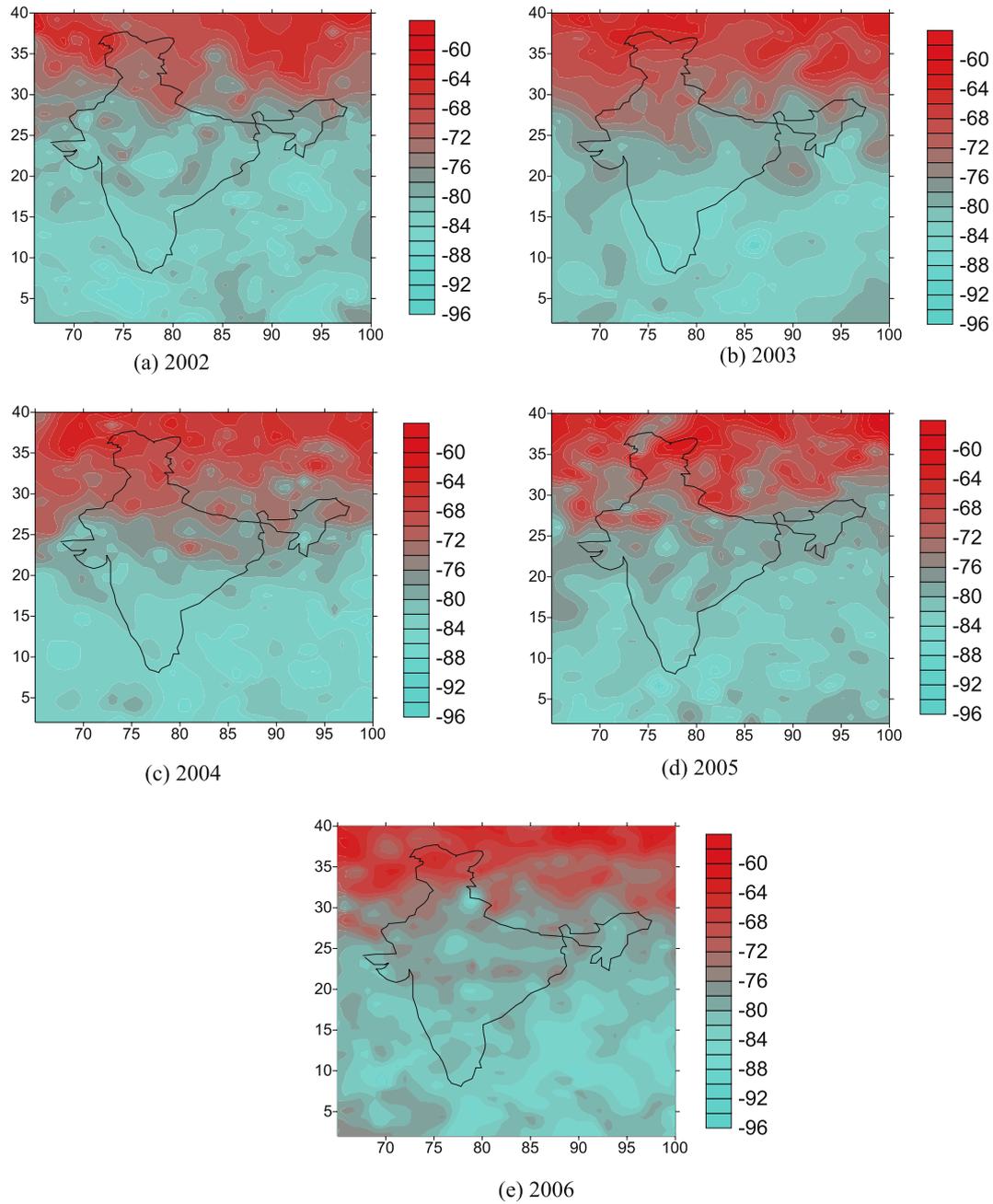


Figure 5. Geographical variation of tropopause temperature in ($^{\circ}\text{C}$) for different years.

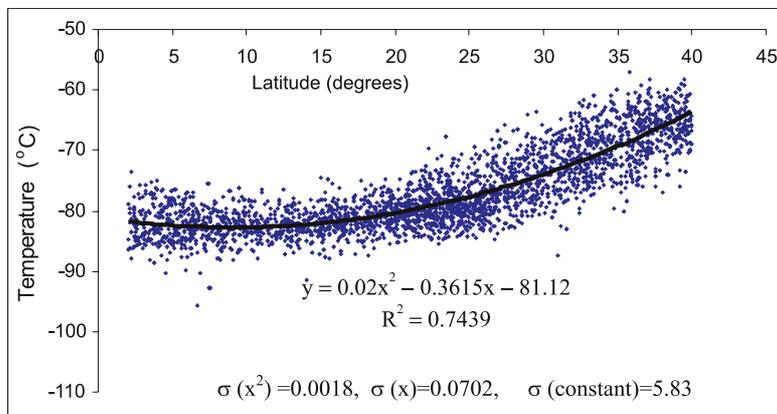


Figure 6. Meridional variation of tropopause temperature.

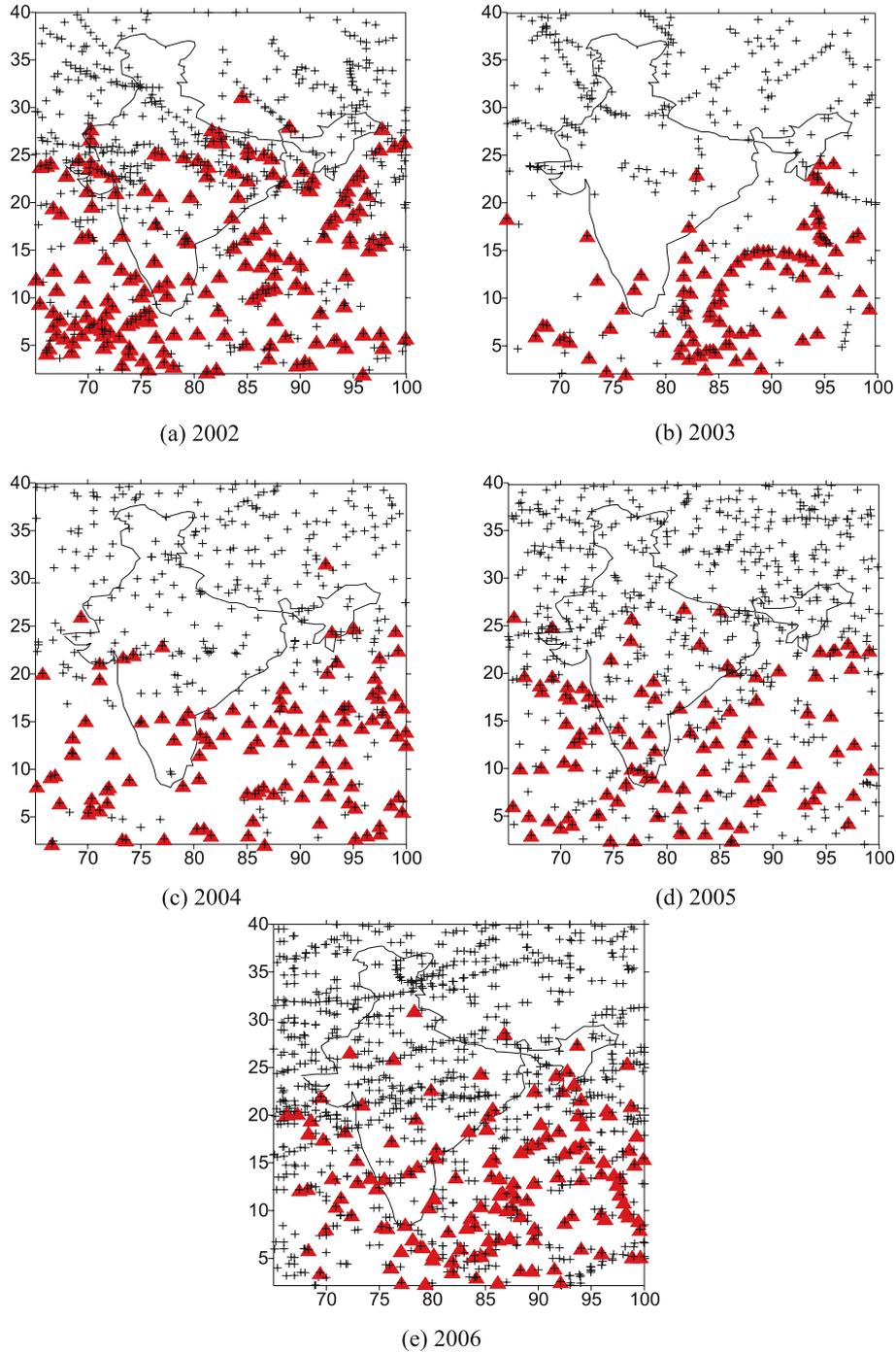


Figure 7. Yearwise distribution of cold tropopause. \blacktriangle represents observation sites where temperature is less than -82°C . $+$ represents observation sites.

(Reid and Gage 1996). Monthly occurrences of cold tropopause for different years are given in table 1. These observations are not in agreement with observations of Jain *et al* (2006); of occurrences of cold tropopause more near to midnight hours and in oceanic region than coastal stations. The tropical upper troposphere is found to be influenced by convective activities occurring in the region and wave forcing in the extra tropics (Reid and Gage

1996). It is reported that cold point temperature depends on the concentration of ozone, carbon dioxide and water vapour in the upper troposphere (Thuburn and Craig 2002).

8. Tropopause in subtropics

It is found that the tropopause shows large variations in the latitude range 30° – 40°N . In this

Table 1. Monthly distribution of tropopause temperature less than -82°C for different years.

Month	Number of occurrences of tropopause temperature less than -82°C				
	2002	2003	2004	2005	2006
January	17	12	18	14	–
February	9	12	15	29	13
March	23	5	25	13	30
April	15	19	18	14	13
May	28	19	1	3	38
June	4	6	–	–	9
July	24	–	–	9	–
August	11	Data not available	–	6	2
September	6	Data not available	–	2	5
October	8	Data not available	–	5	1
November	26	Data not available	3	5	9
December	17	Data not available	31	7	21

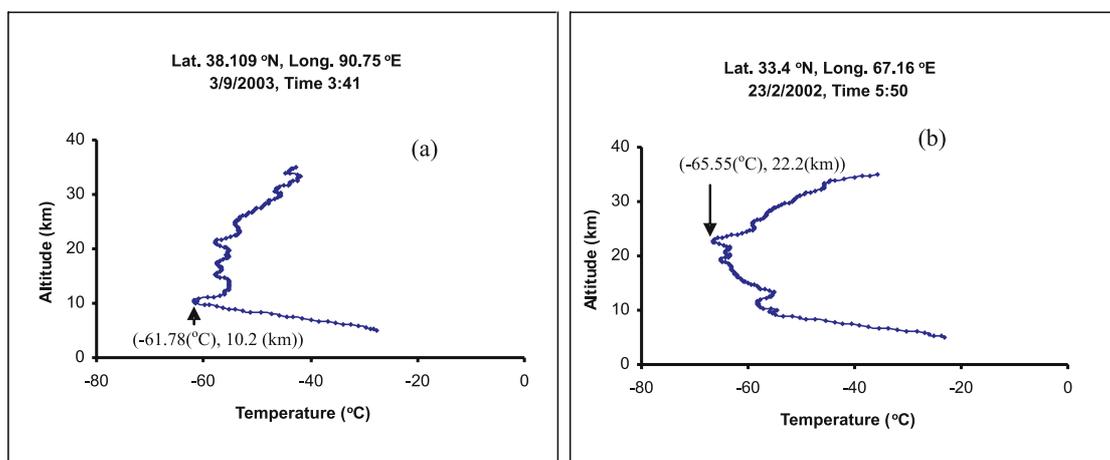


Figure 8. Temperature profiles showing tropopause height and tropopause temperature.

latitude range tropopause height varies from 9 to 23 km. Figure 8 shows some characteristic temperature profiles with tropopause height variation. It is reported on the basis of radiosonde observations in subtropical region that tropopause exhibits bimodal distribution having the characteristics of both tropics and polar region (Seidel and Randel 2007). Tropopause height above 20 km is obtained in 5% of observations made in the latitude range 30° – 40°N . The tropopause heights above 20 km have been reported in some other studies based on lapse rate criteria (Randel *et al* 2007). The distribution of tropopause height with latitude is shown in figure 9. This extremely high tropopause is observed in the months from November to April coinciding with winter months. The tropopause heights below 13 km are also observed in winter months. There is no longitudinal bias obtained in these observations. The tropopause temperature corresponds to extremely high and low tropopause

height, comes in the range from -57° to -73°C . As air parcels are pulled upwards adiabatic, cooling produces temperature below their radiative values and tropopause height increases. While air parcels are pushed downward adiabatic heating produces temperature above their radiative values resulting in lower tropopause heights (Holton *et al* 1995). So these extremely high and low tropopause heights can be considered as an evidence for exchange of airmass between troposphere and stratosphere in this region. The wave induced forces are assumed to exert non-local control over tropospheric stratospheric exchange in extra tropics. The strongest, wave induced forces are occurring in the northern hemisphere during winter season (Holton *et al* 1995). It is reported on the basis of ozone observations that multiple tropopauses observed in this region is associated with deep cyclonic circulation systems of tropical tropospheric air above subtropical jet stream

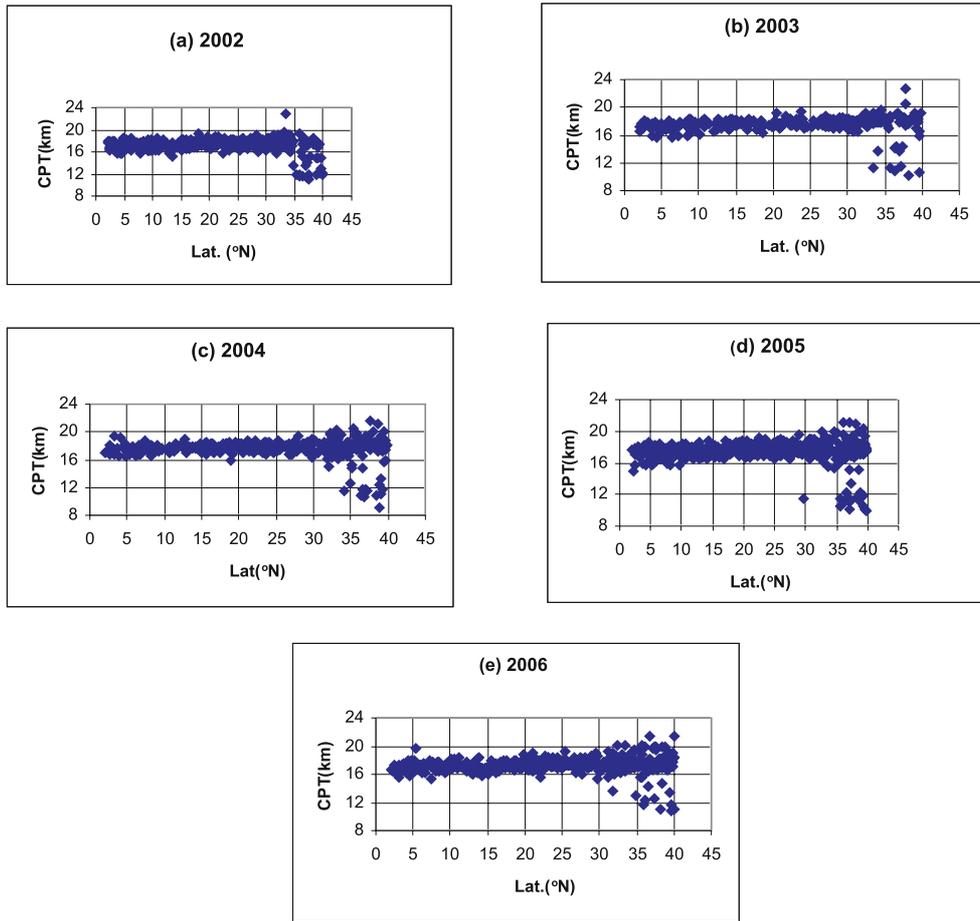


Figure 9. Variation of tropopause height with latitude for different years.

reaching up to lower stratosphere. Extremely high tropopause may be due to some cyclonic circulation systems and transport of tropospheric air into stratosphere as observed in the case of multiple tropopauses. Extremely low tropopause height may be associated with tropopause folds occurring with jet streams and transport of stratospheric air into troposphere. But the occurrence of tropopause fold does not show any seasonality. In our observations, extremely low tropopause heights have been observed only in the winter season. During Rossby wave breaking events near tropopause, stratospheric air enters in the troposphere and tropospheric air enters in stratosphere (Gregory Postel and Mathew Hitchman 2001). These events, also reported in summer months are not coinciding with present observations (Randel *et al* 2007). Further investigation is required to find out exact mechanism behind these atmospheric observations.

9. Conclusion

In the present study, some characteristics of tropopause structure parameters and comparative

study of radiosonde and RO observations over the Indian region have been presented. The utility of RO techniques to study tropospheric climatology is examined. But it is found that in lower troposphere and beyond tropopause region radiosonde observations and RO observations do not agree. In studying weather parameters in the lower troposphere of the region using radio occultation techniques, water vapour climatology suiting to the region and the improved data retrieval techniques are required.

Existence of cold tropopauses indicates that tropical Indian region acts as a source of troposphere to stratosphere transport of airmass. There were studies which suggested that contribution of tropical Indian region to tropospheric to stratospheric transport in monsoon and post monsoon months (Jain *et al* 2006). But present study indicates this transport does not have any seasonal dependency. The large variation of tropopause height in subtropical region suggests the possibility of two-way tropospheric stratospheric exchange of airmass occurring in the region during winter time. The actual mechanism behind such transport has to be studied in detail. In tropical region

tropopause height varies between 15 and 20 km. The existence of high tropopause heights at higher latitudes suggests that the expansion of tropical region is more than that reported by earlier investigators.

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